

Lecture 1.2: Gravity field of the Earth, (T/M Cph. 3).

Gravity direction and magnitude \mathbf{g} (m/s²)

Acceleration

Gravity-potential (W): ability to execute work:

$$Work = \int_A^B |g| ds = W(B) - W(A)$$

The undisturbed sea-surface is a surface where W is constant (we disregard tides etc.)

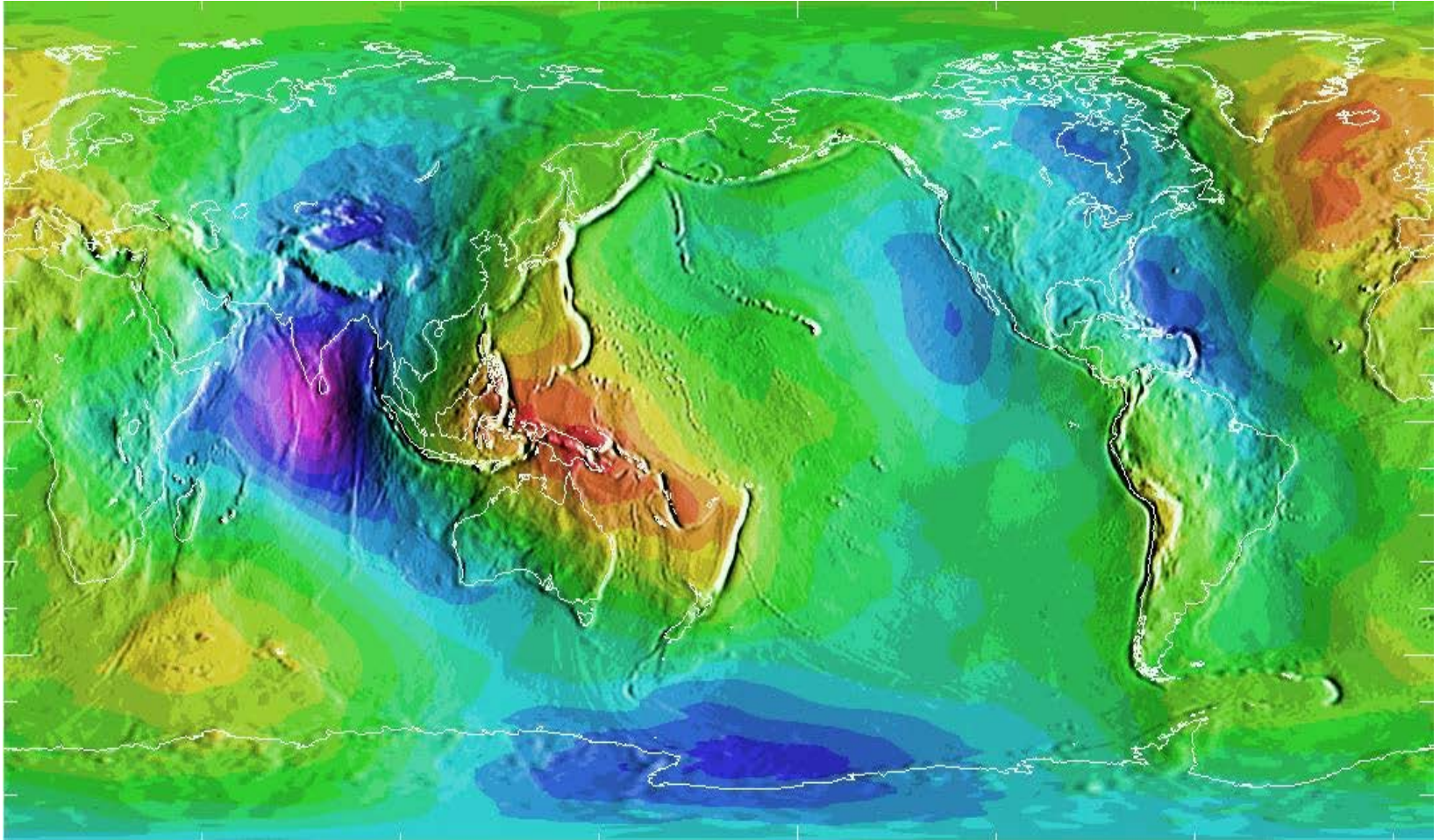
The geoid I

Gravity does not do any work when we move horizontally !

The surface is named the Geoid.

If we can find this, then we know most of the shape of the Earth because most of the Earth is the Ocean.

Global Geoid, min=-107 m, max=86 m



Geoid II

Altitudes, determined by levelling (and gravity measurements) are potential-differences

W and g : Important information on Earth's interior mass-distribution.

Enables levelling with GPS

Important for inertial navigation:

$$\frac{d^2 s}{dt^2} = a - g$$

Units

$$\text{mgal: } 10^{-5} \text{ m/s}^2$$

$$\mu\text{gal: } 10^{-8} \text{ m / s}^2$$

$$\mu\text{ms: } 10^{-6} \text{ m / s}^2$$

$$\text{nms: } 10^{-9} \text{ m / s}^2$$

Spherical coordinates

p. 48

Center in Gravity center

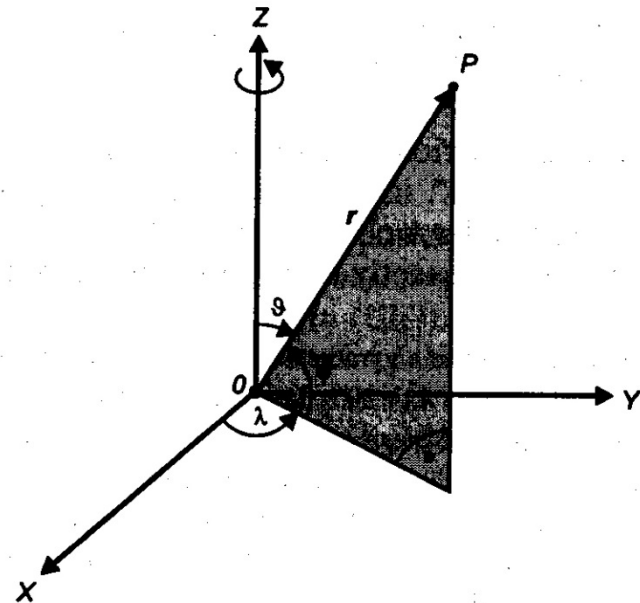
X-Y plane = Equator.

X - axis: Greenwich meridian
Y - axis: 90°
Z - axis: Rotation axis

$$X = r \cos(\bar{\varphi}) \cos(\lambda)$$

$$Y = r \cos(\bar{\varphi}) \sin(\lambda)$$

$$Z = r \sin(\bar{\varphi})$$



Local spherical coordinates

We want to express the gradient of W in a simple way:

$$\left. \begin{array}{l} (1) - \text{axis: East} \\ (2) - \text{axis: North} \\ (3) - \text{axis: Radius - vector} \end{array} \right\} \text{local}(x, y, z)$$

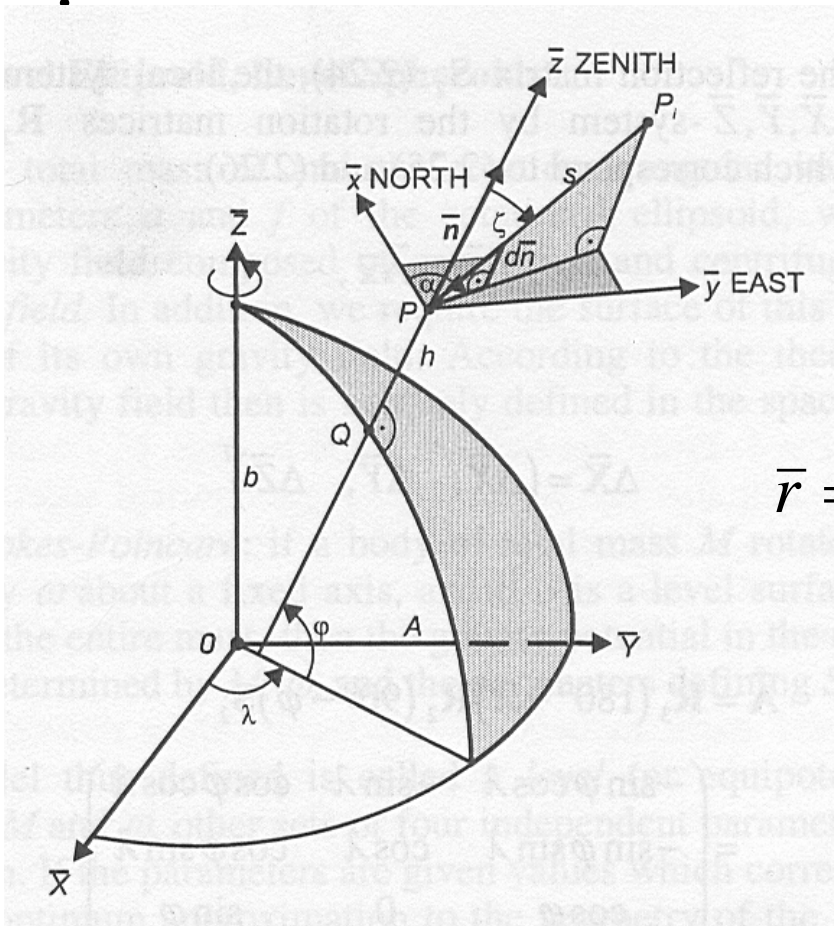
Partial derivatives:

$$\left. \begin{array}{l} \frac{\partial}{\partial x} = \frac{1}{r \cos \bar{\varphi}} \frac{\partial}{\partial \lambda} \\ \frac{\partial}{\partial y} = \frac{1}{r} \frac{\partial}{\partial \bar{\varphi}} \\ \frac{\partial}{\partial z} = \frac{\partial}{\partial r} \end{array} \right\} \text{Proof: use chain - rule}$$

Volume - measure: $d\Omega = r^2 \cos \bar{\varphi} d\lambda d\bar{\varphi}$

Geodetic coordinate-system

p. 98,99



$$\bar{r} = \begin{pmatrix} X \\ Y \\ Z \end{pmatrix} = \begin{pmatrix} (N + h) \cos \varphi \cos \lambda \\ (N + h) \cos \varphi \sin \lambda \\ ((1 - e^2)N + h) \sin \varphi \end{pmatrix}$$

$$N = \frac{a}{(1 - e^2 \sin^2 \varphi)^{1/2}}$$

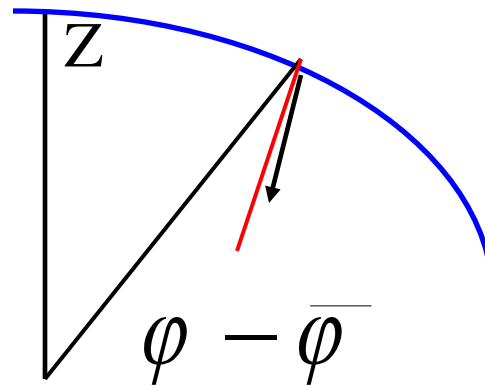
Local geodetic coordinate-system

(1) – axis East

(2) - axis North

(3) - axis: Direction of Ellipsoid-normal

Rotation needed, but small because components of gradientens are small in x and y directions



Newton's law of attraction

p. 54

$$\mathbf{K} = -G \frac{m_1 m_2}{l^2} \left(\frac{\vec{l}}{l} \right)$$

$l = \text{distance (length)},$

$$l = \left((x - x')^2 + (y - y')^2 + (z - z')^2 \right)^{1/2}$$

$\vec{l} = \mathbf{r} - \mathbf{r}' = \text{connection - vector}$

$$G = 6.67259 \times 10^{-11} \text{ m}^3 \text{ kg}^{-1} \text{ s}^{-2}$$

Gravity acceleration.

Gravity acceleration: $m_2=1$

$$\mathbf{b} = -G \frac{\mathbf{r} - \mathbf{r}'}{|\mathbf{r} - \mathbf{r}'|^3} \quad \text{then}$$

$$\text{curl} \mathbf{b} = \frac{\partial}{\partial x} b_1 + \frac{\partial}{\partial y} b_2 + \frac{\partial}{\partial z} b_3 = 0$$

$\Rightarrow \mathbf{b}$ gradient of potential

$$V = \frac{Gm}{l} \quad (\text{sign geophysical convention})$$

For the Earth

$$V = G \iiint_{Earth} \frac{dm}{l}$$
$$= G \iiint_{Earth} \frac{\rho}{l} dv, \quad \rho = \textit{massedensity}$$

Special case: Spherical symmetric Earth. p. 55

Shell, σ : thickness dr' , density ρ , constant

$$V' = G\rho dr' \iint_{\sigma} \frac{1}{l} (r')^2 \cos \bar{\varphi} d\bar{\varphi} d\lambda,$$

For exterior potential, $r > r'$

$$V' = 4\pi G\rho (r')^2 dr' \frac{1}{r},$$

then $dm' = 4\pi\rho (r')^2 dr'$ (mass of shell)

$$\text{Exterior potential } V_e = \frac{GM}{r}$$

Potential inside.

P. 56,57

Interior potential, constant density T(3.21)

$$\begin{aligned} V_i &= \frac{4}{3} \pi G r^2 \int_0^r \rho r'^2 dr' + 4 \pi G \int_r^R \rho r' dr' \\ &= 2 \pi G \rho \left(R^2 - \frac{r^2}{3} \right) \end{aligned}$$

Gradient of V.

P.58

$$\frac{\partial \mathcal{V}}{\partial x} = V_x = -G \iiint_{Earth} \frac{x - x'}{l^3} dm$$

$$\frac{\partial^2 V}{\partial x^2} = -G \iiint_{Earth} \frac{1}{l^3} dm + 3G \iiint_{Earth} \frac{(x - x')^2}{l^5} dm$$

$$\Rightarrow \Delta V = \sum_{i=1}^3 \frac{\partial^2 V}{\partial x_i^2} = V_{xx} + V_{yy} + V_{zz} = 0$$

V is HARMONIC function outside masses

Inside: $\Delta V = -4\pi G\rho$, POISSONS equation

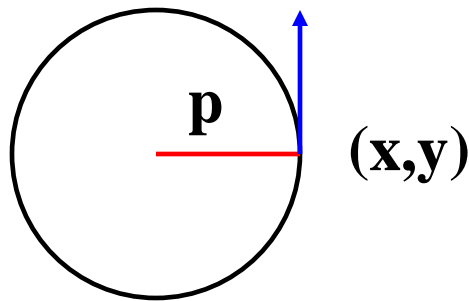
Centrifugal-force

p. 60

We rotate in orbit which is a circle around the Z-axis:

$$x(t) = p \times \cos(\omega(t - t_0))$$

$$y(t) = p \times \sin(\omega(t - t_0)), \quad z = \textit{constant}$$



Centrifugal potential.

Velocity:

$$\dot{x} = -p \times \omega \times \sin(\omega(t - t_0))$$

$$\dot{y} = p \times \omega \times \cos(\omega(t - t_0)), \quad \dot{z} = 0$$

Acceleration:

$$\ddot{x} = -p \times \omega^2 \cos(\omega(t - t_0))$$

$$\ddot{y} = -p \times \omega^2 \sin(\omega(t - t_0)), \quad \ddot{z} = 0.$$

Gradient of:

$$\Phi(x, y, z) = \frac{\omega^2}{2} (x^2 + y^2) = \frac{\omega^2}{2} r^2 \cos^2 \bar{\varphi}$$

Total potential.

p. 62

$$W = V + \Phi = G \iiint_{Earth} \frac{\rho}{l} dv + \frac{\omega}{2} p^2$$

$$\mathbf{g} = \nabla W,$$

$$|\mathbf{g}| = \textit{gravity}$$

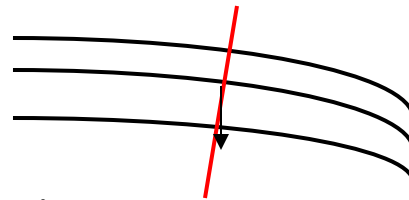
Generalised Poisson Equation

$$\Delta \mathbf{W} = -4\pi \mathbf{G} \rho + 2\omega^2$$

Level surfaces.

p. 63

W constant, cut by plumb-line \mathbf{g} tangent.



Plumb-line coordinate axis, so

dW is a total differential, $dW = -gdn + 0dx + 0dy$

Height difference between 2 points is **independent** of the path we take between them. Used to adjust levelling.